

1 BACKGROUND

What is the change in height by the densification of firn on glaciers and ice caps?

Geodetic elevation change measurements are a promising tool to derive high resolution mass balances of glaciers and ice caps. Up to now, geodetic mass balances inherit an uncertainty which occurs when converting geodetic volume change into glacier mass balance. The inaccuracy results from an inadequate conversion factor (density) which neglects altitude dependent firn density variations, firn layer thickness and not homogenous density variations with varying climate conditions.

To improve the accuracy of glacier mass balances measured in geodetic manner we develop a transferable firn elevation change model to minimize this systematic error. The main scope of the firn model are short term GMBs (3-5 yr) due to a high variability in the possible conversion factor^[13].

3 RESULTS

Exemplary model run:

We selected four MAR modelpoints (fig. a) for a 9 year test phase, with 20 m depth, 0.5 m layer thickness, an idealized bottom temperature of -3.7°C^[7] and one year spin up. Fig. b gives a brief overview about firn elevation change in meter. The firn evolution fits to the homogeneous snowpack pattern over the ice cap reported by Möller et al. (2011)^[3].

Geodetic measurements:

The second project part comprises an extensive analysis of the geodetic mass balance of the test site. Fig. d & e give a first impression on how the elevation of two TanDEM-X DEMs compared to one ICESat track (spring 2004) has changed between 2004 and 2012. The firn compaction signal is not yet removed, nevertheless, mind the thickening on the northwestern face of the ice cap (fig. e).

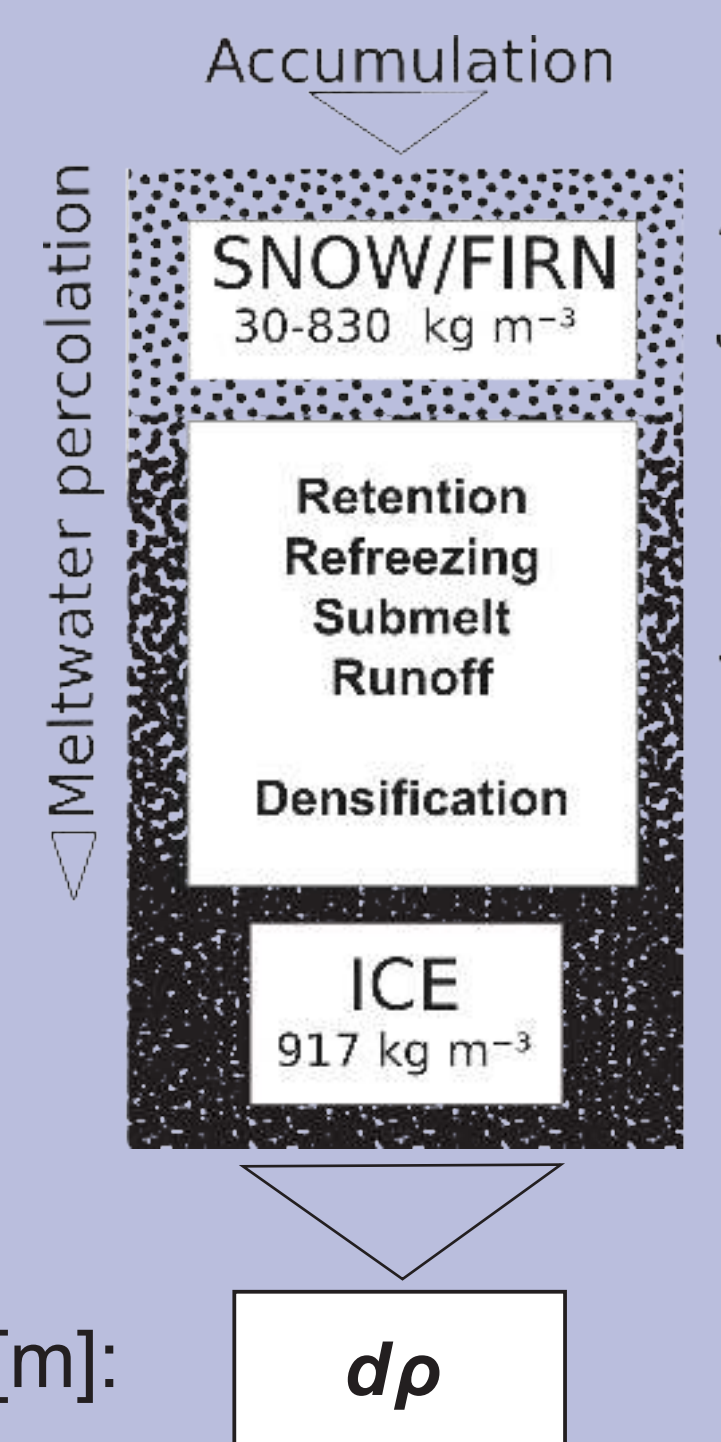
2 MODEL

The model is implemented in Python 2.7 and is based on a surface energy and snowpack model (COSIMA^[5]).

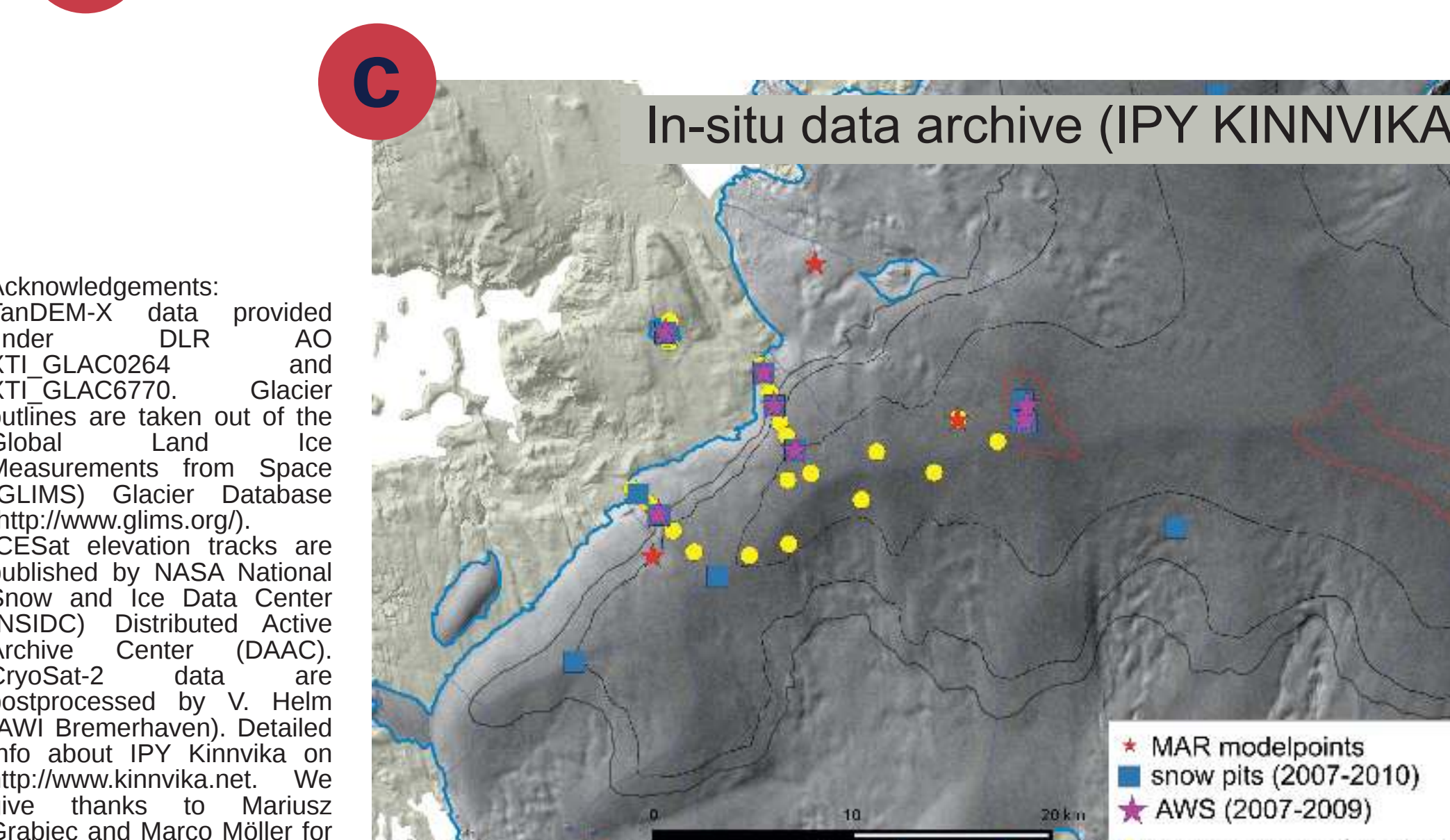
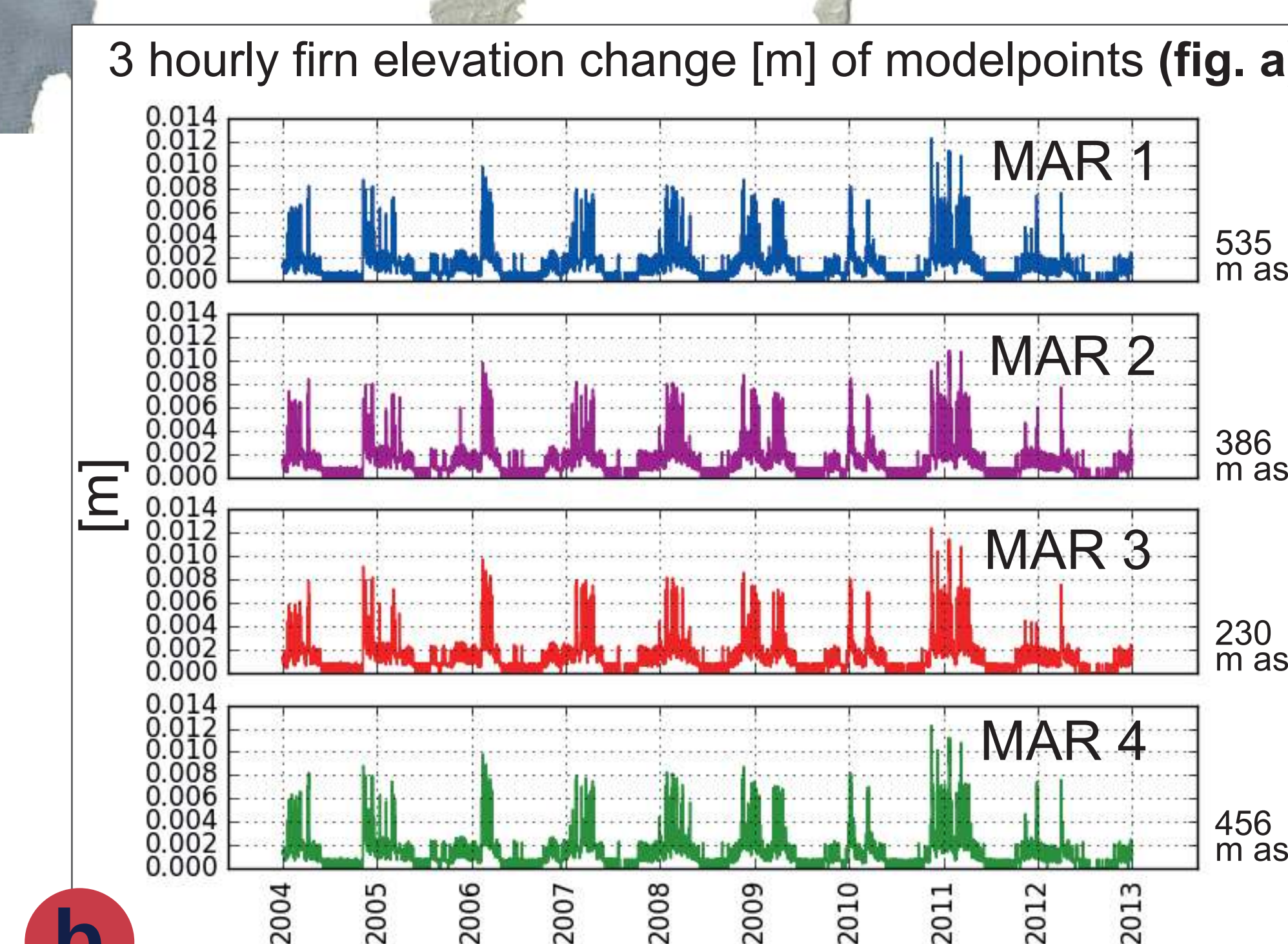
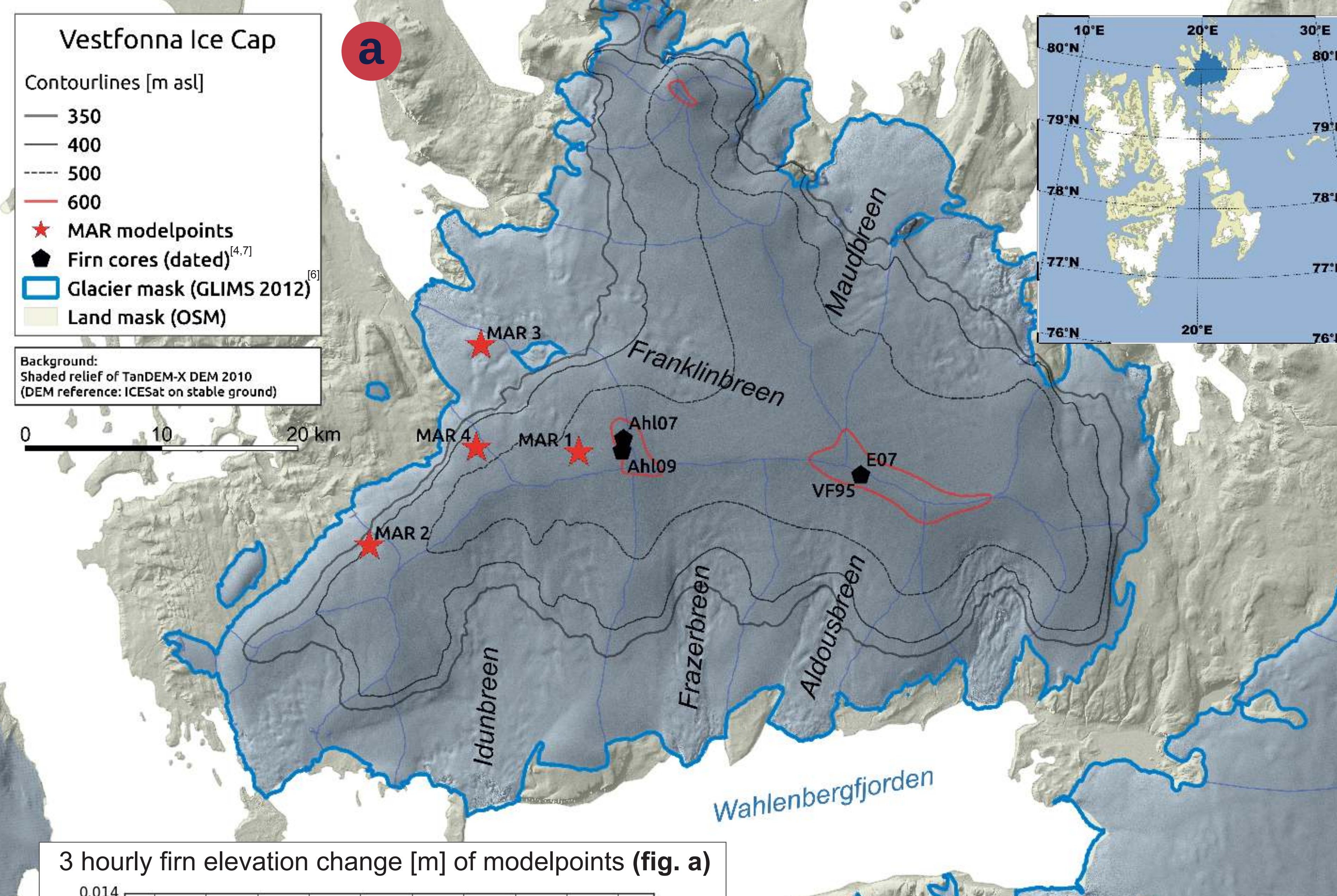
Forcing: subdaily reanalysis (MAR^[11]) data (fig. a):

- air temperature [K]
- incoming shortwave radiation [W m⁻²]
- relative humidity [%]
- air pressure [hPa]
- wind speed [m s⁻¹]
- all-phase precipitation [m]
- cloud cover fraction.

Model concept:



$$dM/dt = (dh/dt - dp/dt) * S * \rho_{ice}$$



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4 The test site

Vestfonna Ice Cap (VIC) (~2340 km²) on Nordaustlandet, north east of Svalbard.

- Vertical spanning: 0 - 630 m asl
- Mean elevation: 386 m asl
- Mean ELA: 380 m asl (S↔N)^[2]

The VIC is a polythermal ice cap with a dome like shape and gentle slopes. Most outlet glaciers calve into surrounding seas. Melting can occur up to summit in summer season and mass balance year starts in September, with ablation period from June to August.^[2]

The VIC showed almost balanced conditions in the last two decades, while the outlets were steady retreating, with the exception of the re-advance of the largest outlet glacier towards NW (surge).^[2,10,11]

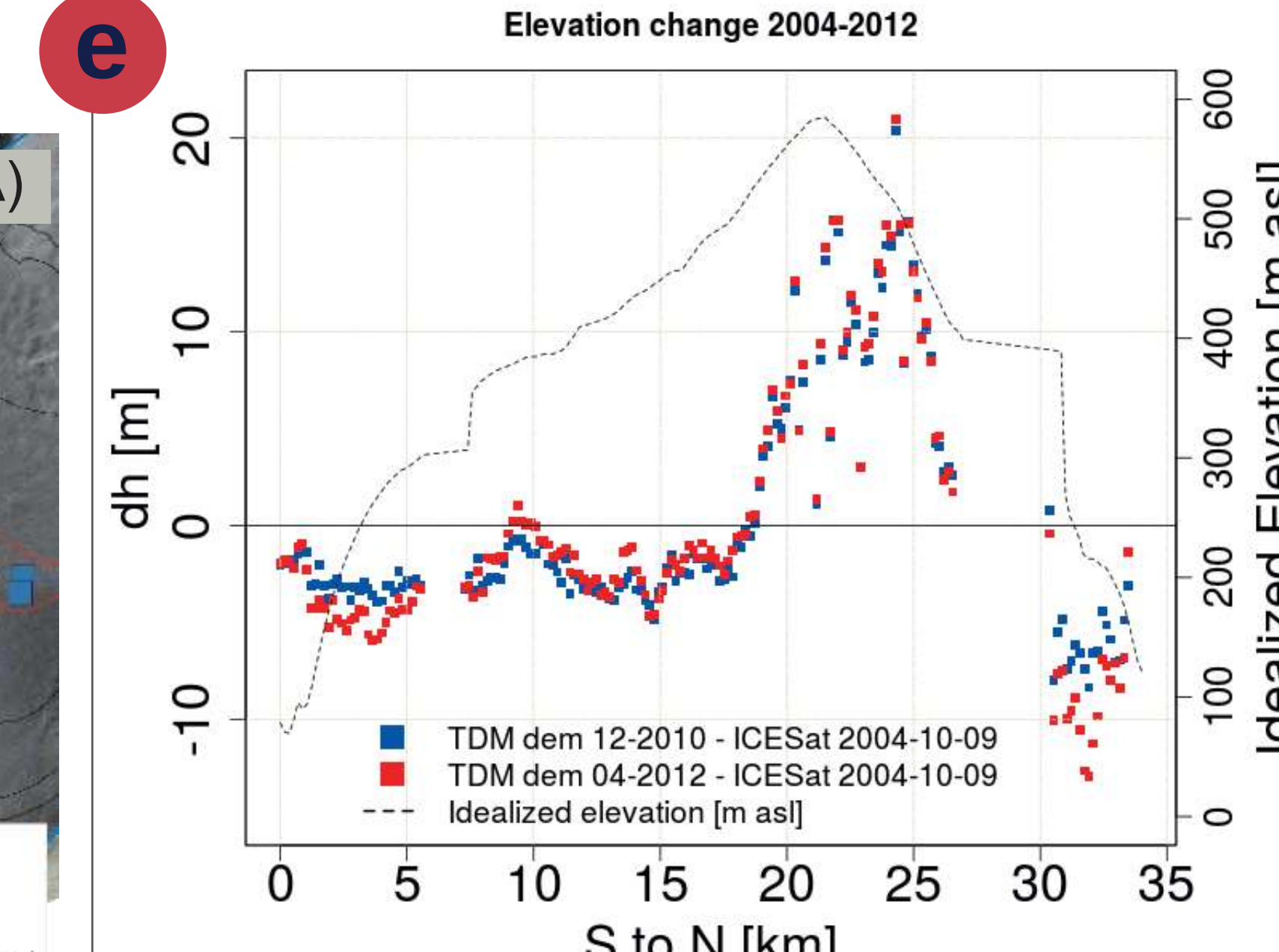
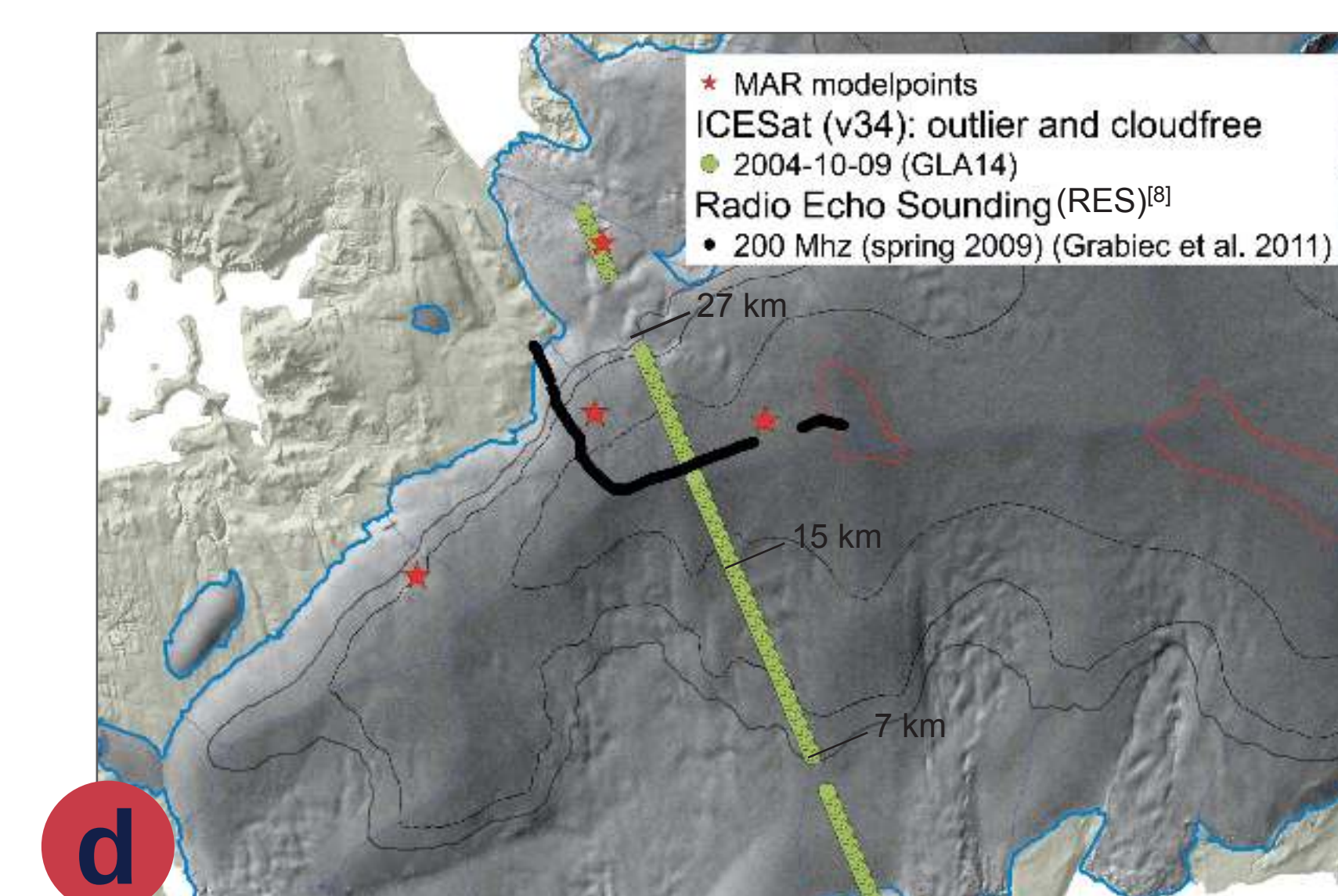
Firn thickness was 1996 10 m at summit^[7], recent studies suggest it to be 15 - 20 m thick^[4]. Melt water percolation and refreezing play a major role on the ice cap.

5 CONCLUSION

■ For calibration and validation we have to run sensitivity tests against in-situ measurements (stakes (SMB)^[2], snow pits (density & depth)^[3], RES (stratigraphy)^[8], firn cores (aging)^[4]) (fig. c) to tune the parametrization.

■ The elevation gain in the interior of VIC approve the observed accumulation increase in Northeastern Svalbard and a corresponding elevation increase uphill^[11,14].

■ TDM DEMs for Dec 2010 and April 2012 are ready, but they show no significant elevation change (signal is in error bar)! TDM DEM for 2015 is in process.



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